

and hence

$$\frac{\partial F^*}{\partial t} = \frac{\partial F}{\partial t} + \frac{\partial F}{\partial \theta} \frac{\partial \theta}{\partial t}, \quad (4-174)$$

$$\left. \frac{\partial F}{\partial t} \right)_{\Theta=\text{const.}} = \left. \frac{\partial F}{\partial t} \right)_{\theta=\text{const.}} + F_{\theta} \frac{\partial \theta}{\partial t}, \quad (4-175)$$

in an obvious notation. Thus, in order to get  $\partial F/\partial t$  in Wavre's sense, we have to add to  $\partial F/\partial t$  in our present sense a " $\theta$ -correction".

The factor  $\partial \theta/\partial t$  is the change of  $\theta$  along the normal to the equisurface passing through the point  $(t, \theta)$  under consideration. It is easily found as follows (Fig. 4.9). The infinitesimal distance  $PF$  can be expressed in two ways:

$$-r d\theta = \delta dr \quad (4-176)$$

(we have put the minus sign since in Fig. 4.8 we had taken  $r = OP_1$ , whereas now

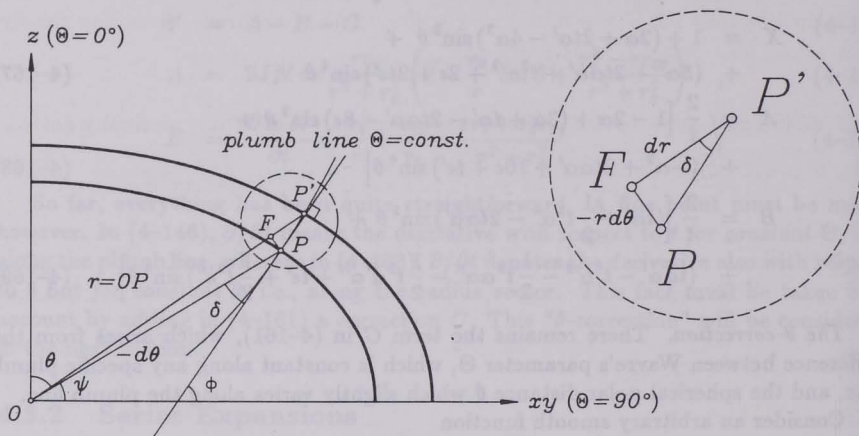


FIGURE 4.9: The  $\theta$ -correction

$r = OP$ ; so to speak, in Fig. 4.8 we went from  $P^i$  to  $P$ , whereas in Fig. 4.9 we go from  $P$  to  $P^i$ ). Thus

$$\frac{\partial \theta}{\partial r} = -\frac{\delta}{r}, \quad (4-177)$$

where the very small angle  $\delta$  is nothing else than the difference between the geographic latitude  $\phi$  and the geocentric latitude  $\psi$  (Fig. 4.9), which is given by (1-76):

$$\delta = \phi - \psi = 2f \cos \theta \sin \theta, \quad (4-178)$$

neglecting higher-order terms. (This is a standard formula from ellipsoidal geometry: to this accuracy, the level surfaces can be considered ellipsoids of revolution.) To the